
Introduction to the
Spectral-infinite-element method

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Outline

- Background
- Infinite-element method
- Spectral-element method
- Examples

Global problems

- Wave propagation
- Postearthquake relaxation
- Glacial isostatic adjustment
- Tidal loading



Governing equations

- Conservation of linear momentum

$$\nabla \cdot \boldsymbol{\sigma} + \nabla(\rho \mathbf{u} \cdot \mathbf{g}) - \nabla \cdot (\rho \mathbf{u}) \mathbf{g} - \rho \nabla \phi = \rho \ddot{\mathbf{u}} + 2\rho \boldsymbol{\Omega} \times \dot{\mathbf{u}} + \mathbf{f}$$

- Conservation of angular momentum

$$\dot{\mathbf{H}} + \boldsymbol{\Omega} \times \mathbf{H} = \mathbf{0}$$

- Gravity potential

$$\nabla^2 \phi + 4\pi G \nabla \cdot (\rho \mathbf{s}) = 0$$

SPECFEM3D Cartesian

- Conservation of linear momentum

$$\nabla \cdot \boldsymbol{\sigma} + \nabla(\rho \mathbf{u} \cdot \mathbf{g}) - \nabla \cdot (\rho \mathbf{u}) \mathbf{g} - \rho \nabla \phi = \rho \ddot{\mathbf{u}} + 2\rho \boldsymbol{\Omega} \times \dot{\mathbf{u}} + \mathbf{f}$$

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SPECFEM3D Globe

- Conservation of linear momentum

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$$\dot{\mathbf{H}} + \boldsymbol{\Omega} \times \mathbf{H} = \mathbf{0}$$

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Motivation

- Conservation of linear momentum

$$\nabla \cdot \boldsymbol{\sigma} + \nabla(\rho \mathbf{u} \cdot \mathbf{g}) - \nabla \cdot (\rho \mathbf{u}) \mathbf{g} - \rho \nabla \phi = \rho \ddot{\mathbf{u}} + 2\rho \boldsymbol{\Omega} \times \dot{\mathbf{u}} + \mathbf{f}$$

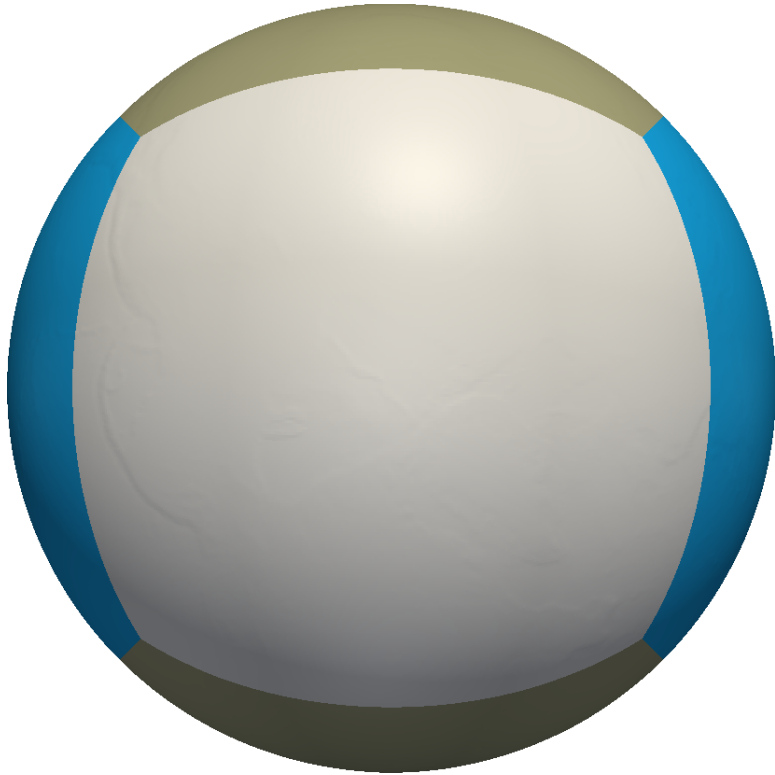
- Conservation of angular momentum

$$\dot{\mathbf{H}} + \boldsymbol{\Omega} \times \mathbf{H} = \mathbf{0}$$

- Gravity potential

$$\nabla^2 \phi + 4\pi G \nabla \cdot (\rho \mathbf{s}) = 0$$

Problem size



- $NEX = 256$
- $N_{spec} = 4,450,304$
- $N_{node} = 299,068,155$

Problem size

Explicit scheme

$$\ddot{\mathbf{s}} = \mathbf{M}^{-1}(\mathbf{f} - \mathbf{C}\dot{\mathbf{s}} - \mathbf{K}\mathbf{s})$$

- $N_{\text{dof}} = 1,196,272,620$
- Maximum array size = 1,196,272,620

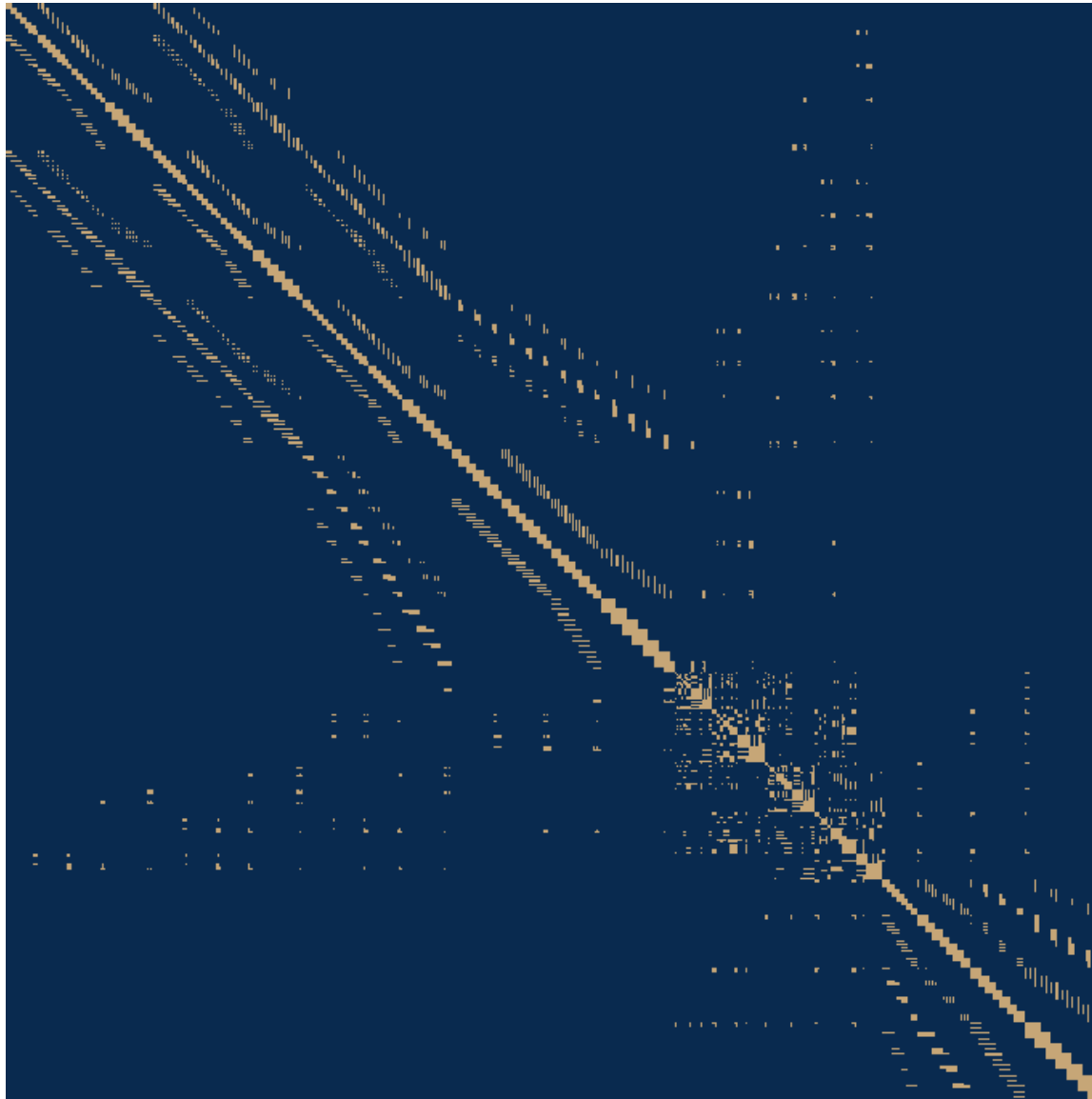
Implicit scheme

$$\mathbf{s} = \mathbf{K}^{-1} \mathbf{f}$$

- $N_{\text{dof}} = 1,196,272,620$
- Maximum array size = 1,196,272,620 X 1,196,272,620

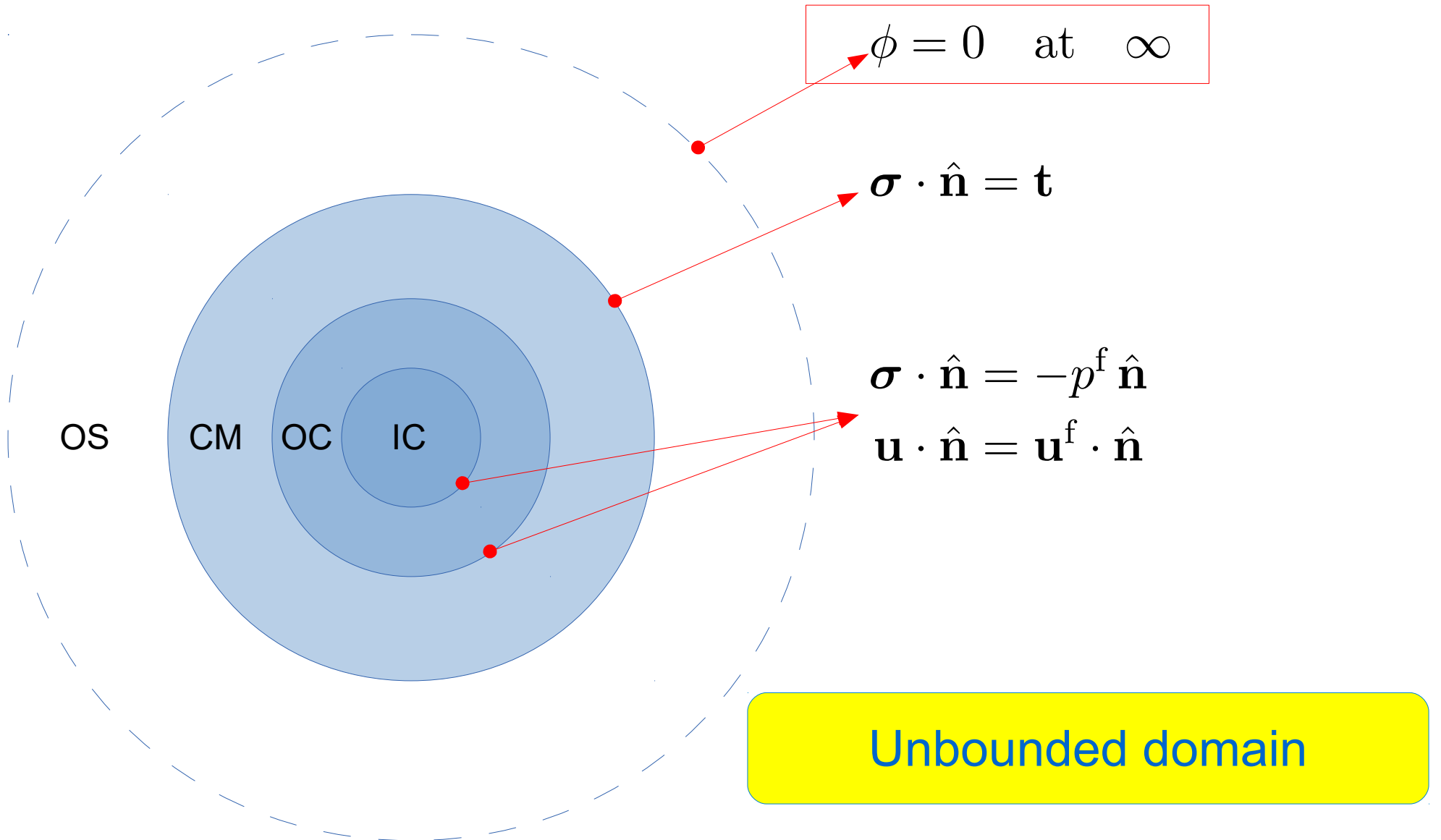
Huge problem!

Problem size



- Sparsity
- Element-by-element

Gravity

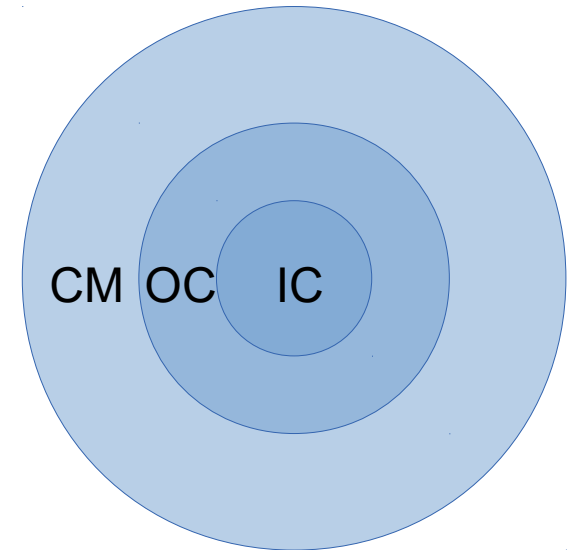
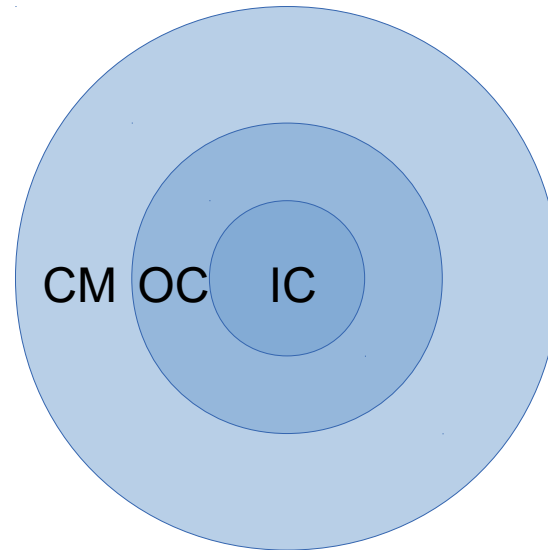
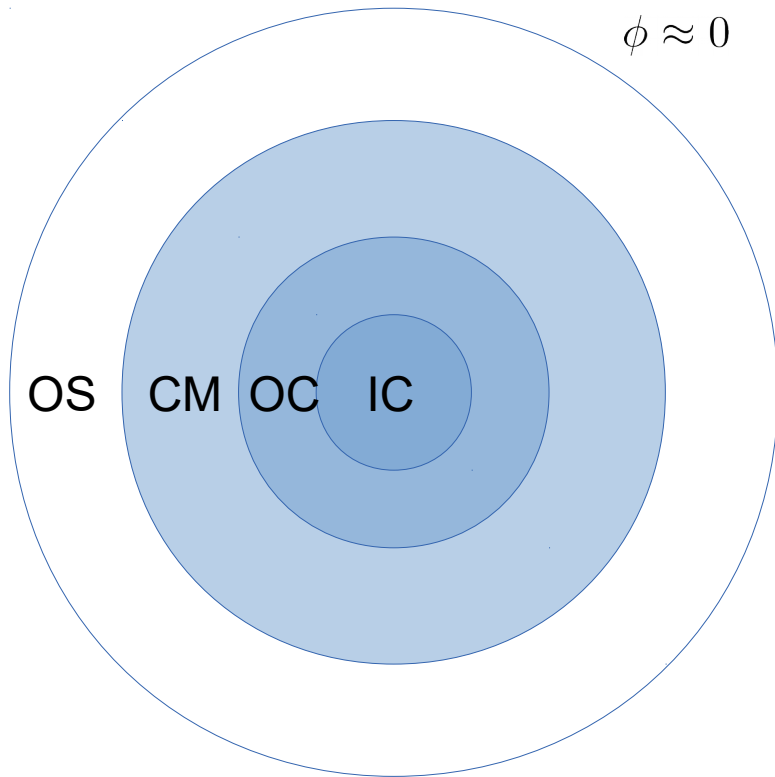


Existing methods

Domain extension

Spherical harmonics

First principle

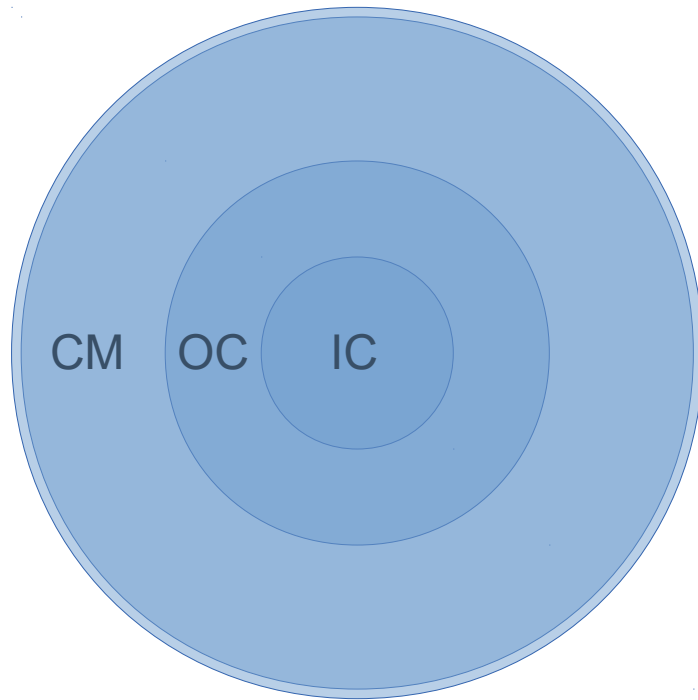


Is it possible to consider full 3D model?

$$\begin{aligned} \phi_s &= -G \int \frac{\rho(\mathbf{r})[\mathbf{s}(\mathbf{r}) \cdot (\mathbf{r}_s - \mathbf{r})]}{|\mathbf{r}_s - \mathbf{r}|^3} d\Omega \\ &- G \int_{r_s=R} \frac{\sigma(\mathbf{r}_s)}{|\mathbf{r}_s - \mathbf{r}|} dS - G \int_{r_s=R_{OC}} \frac{\sigma(\mathbf{r}_s)}{|\mathbf{r}_s - \mathbf{r}|} dS \\ &- G \int_{r_s=R_{IC}} \frac{\sigma(\mathbf{r}_s)}{|\mathbf{r}_s - \mathbf{r}|} dS \end{aligned}$$

Infinite-element method

(Zienkiewicz et al, 1983)



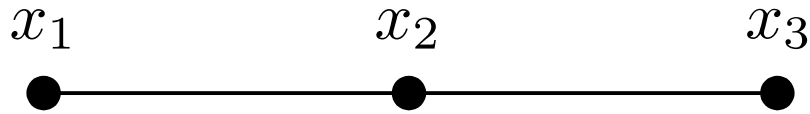
Infinite-element layer

Widely used in solid/fluid mechanics!

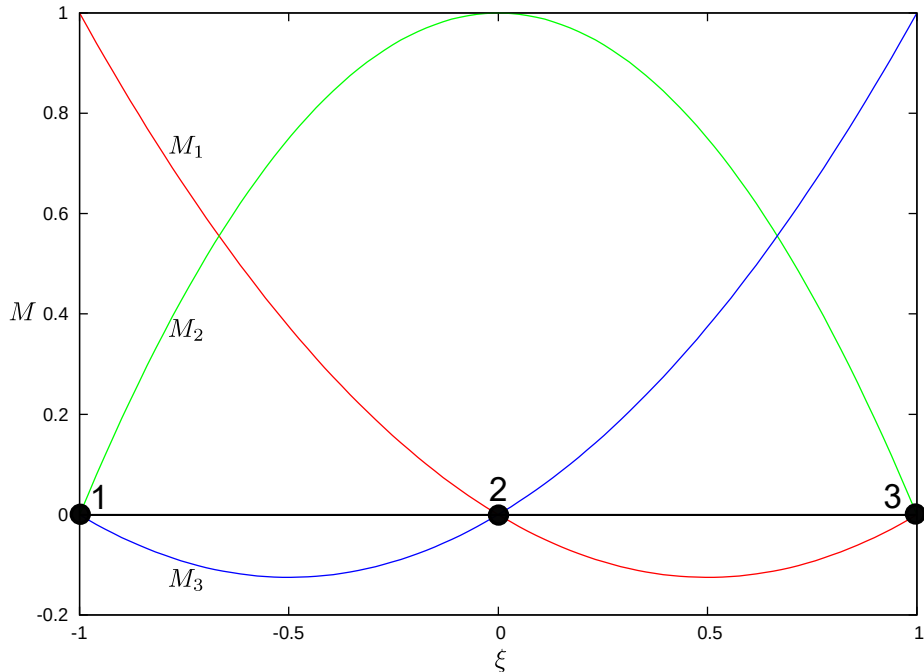
Infinite-element mapping

(Zienkiewicz et al, 1983)

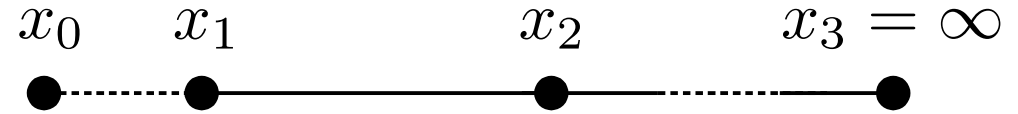
Finite element



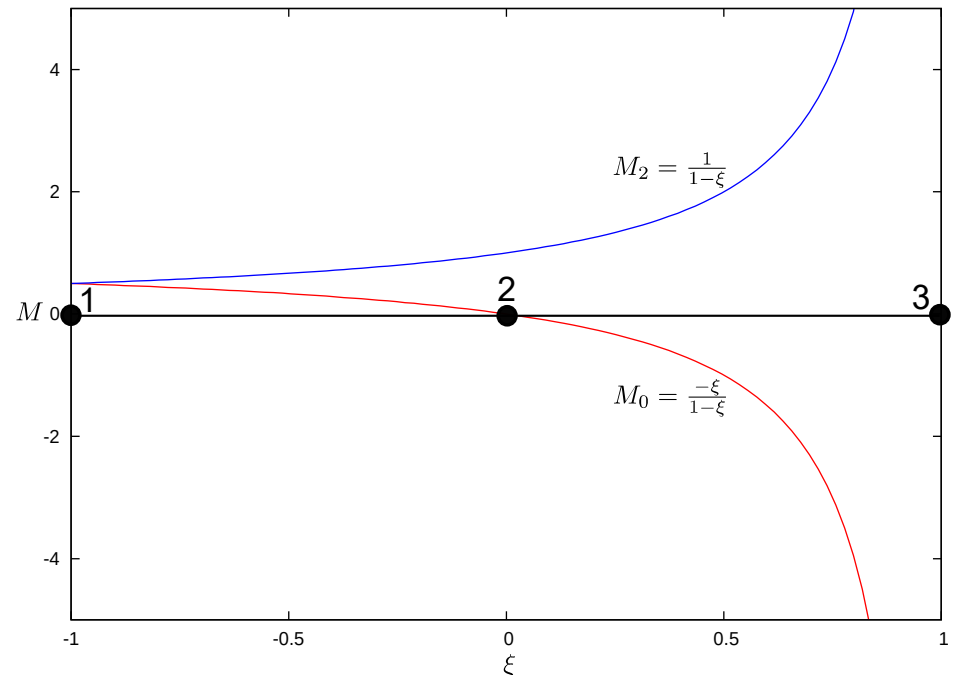
$$x = \sum_{i=1}^3 M_i(\xi) x_i$$



Infinite element



$$x = M_0(\xi) x_0 + M_2(\xi) x_2$$



Why it works?

$$\phi(\xi) = a_0 + a_1\xi + a_2\xi^2 + a_3\xi^3 + \dots$$



$$x = M_0(\xi) x_0 + M_2(\xi) x_2$$

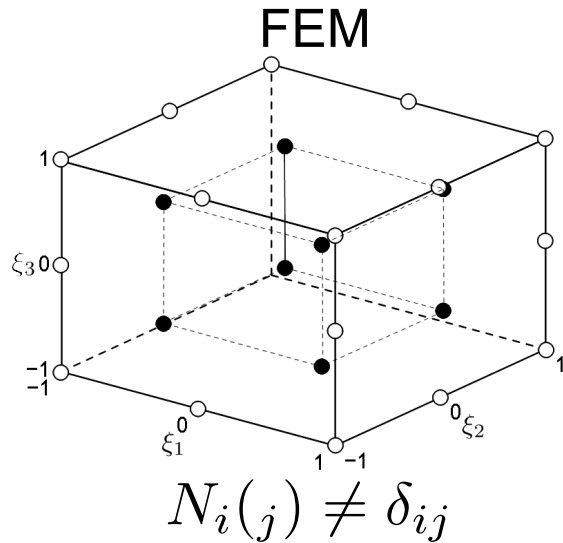
$$r = x - x_0$$

$$\phi(r) = b_0 + \frac{b_1}{r} + \frac{b_2}{r^2} + \frac{b_3}{r^3} + \dots$$

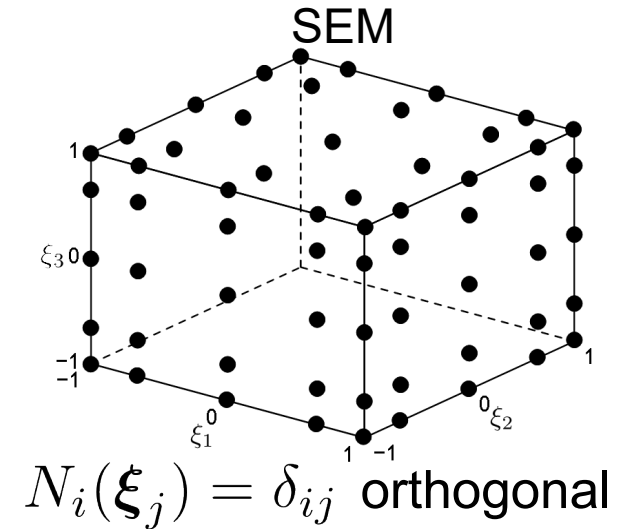
ϕ decays to b_0 as r tends to infinity

Spectral-element method

(e.g., Patera 1984, Cohen et al 1993, Faccioli et al 1997, Komatitsch & Tromp 1999)



- Interpolation
- Integration



$$\begin{bmatrix} * & * & * & * & * \\ * & * & * & * & * \\ * & * & * & * & * \\ * & * & * & * & * \\ * & * & * & * & * \end{bmatrix}$$

Mass matrix

$$\mathbf{M} = \int_{\Omega} \rho N_i N_j d\Omega$$

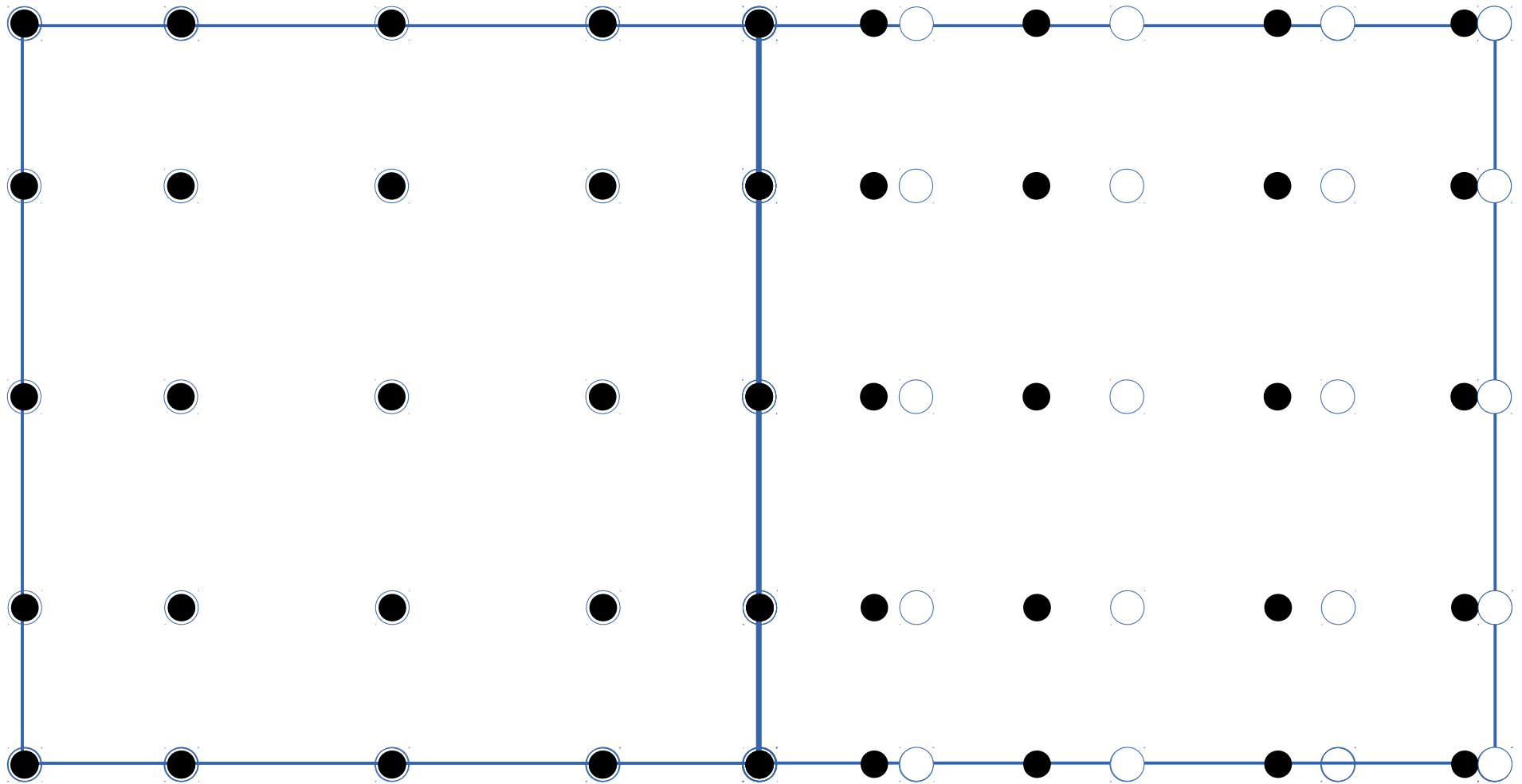
$$\begin{bmatrix} * & & & & \\ & * & & & \\ & & * & & \\ & & & * & \\ & & & & * \end{bmatrix}$$

Flexibility of FEM and stability of spectral method!

Spectral-infinite-elements

Spectral element

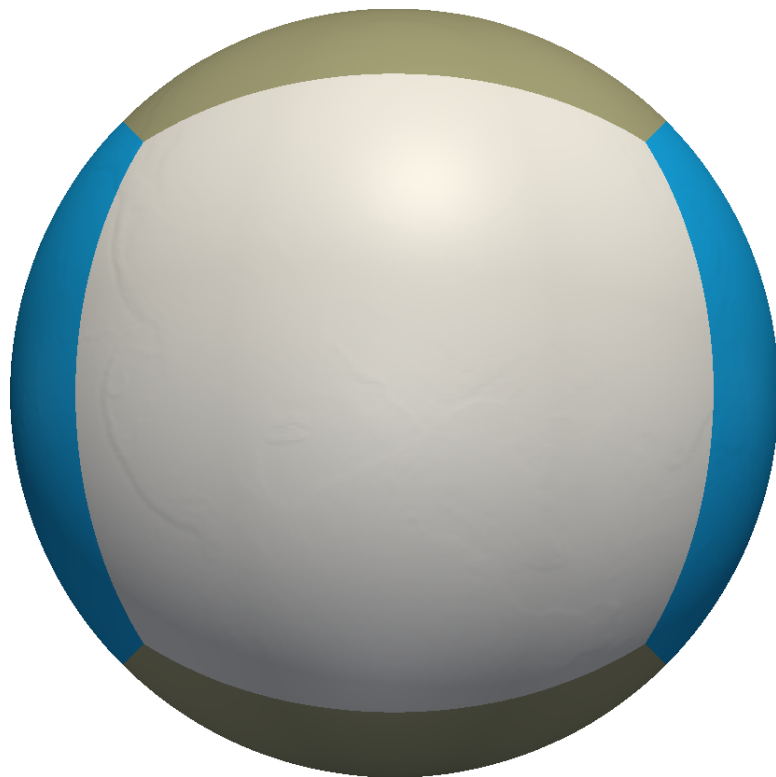
Infinite element



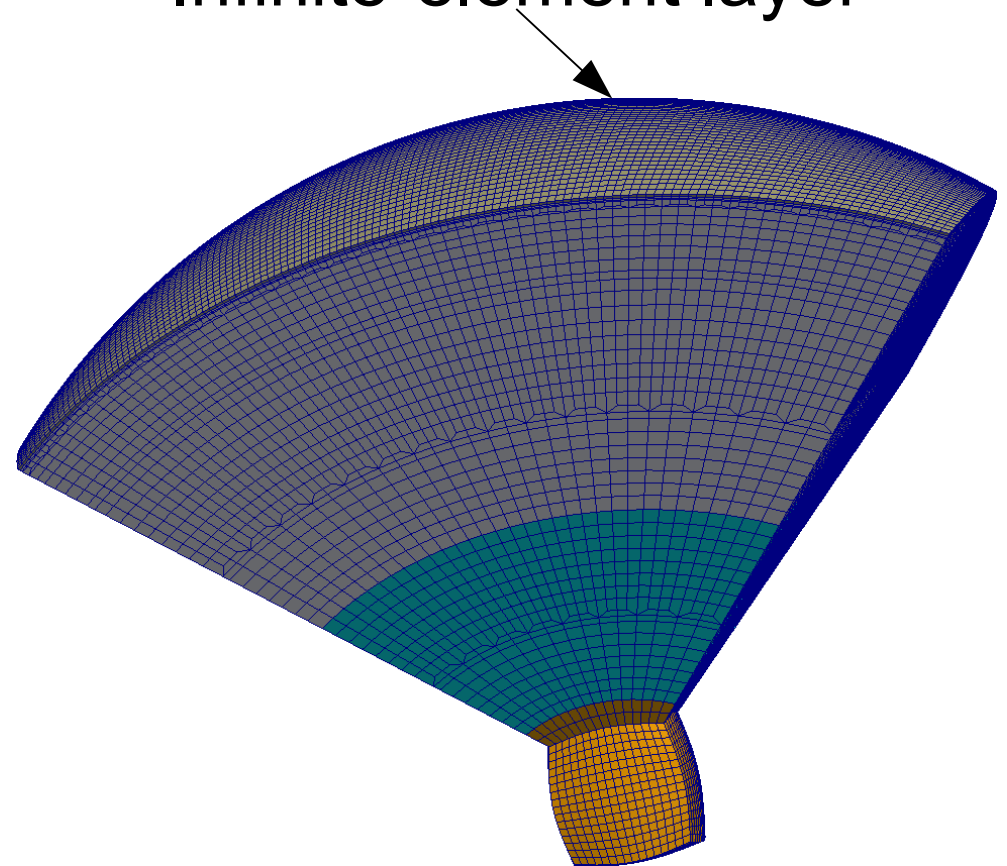
- Interpolation points
- Quadrature points

Lobatto quadrature vs Radau quadrature

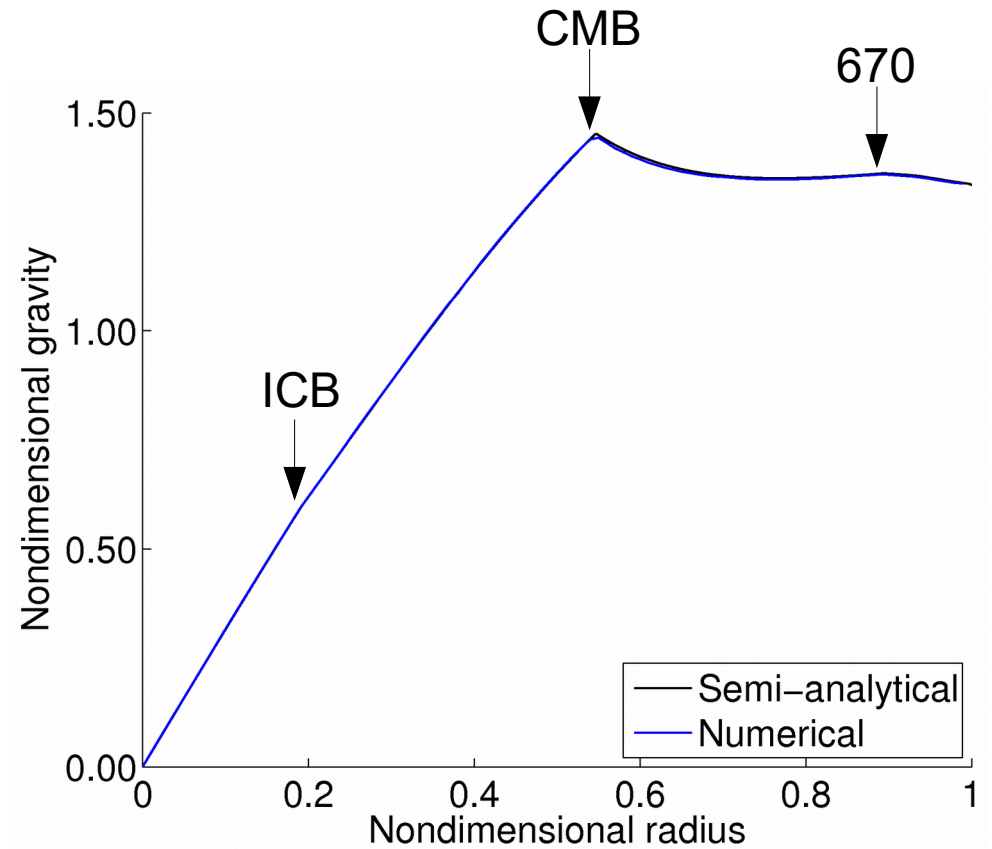
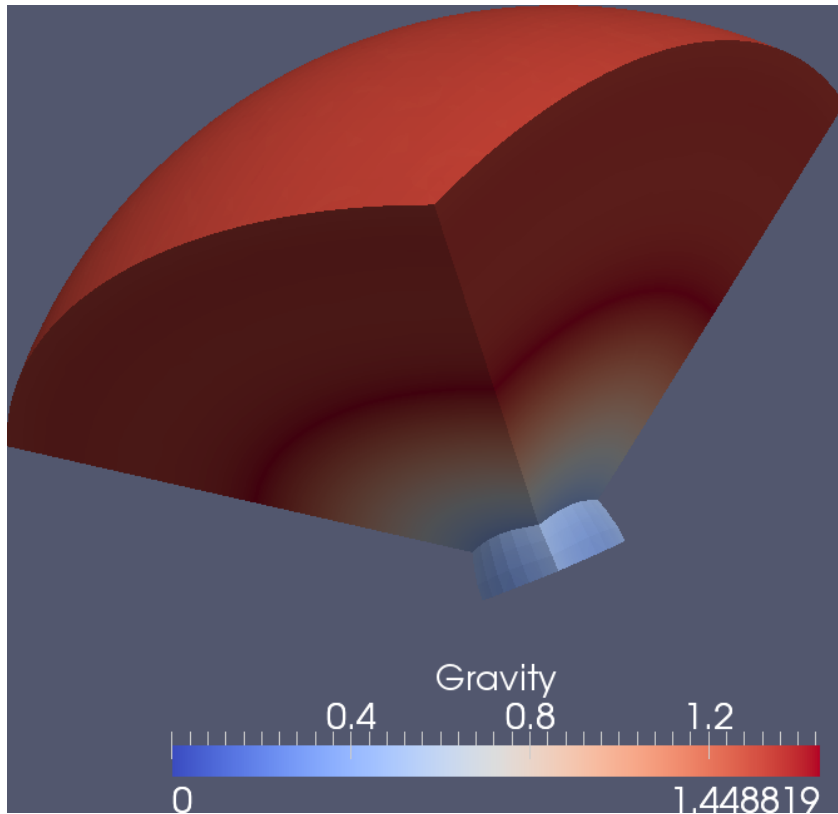
Meshing



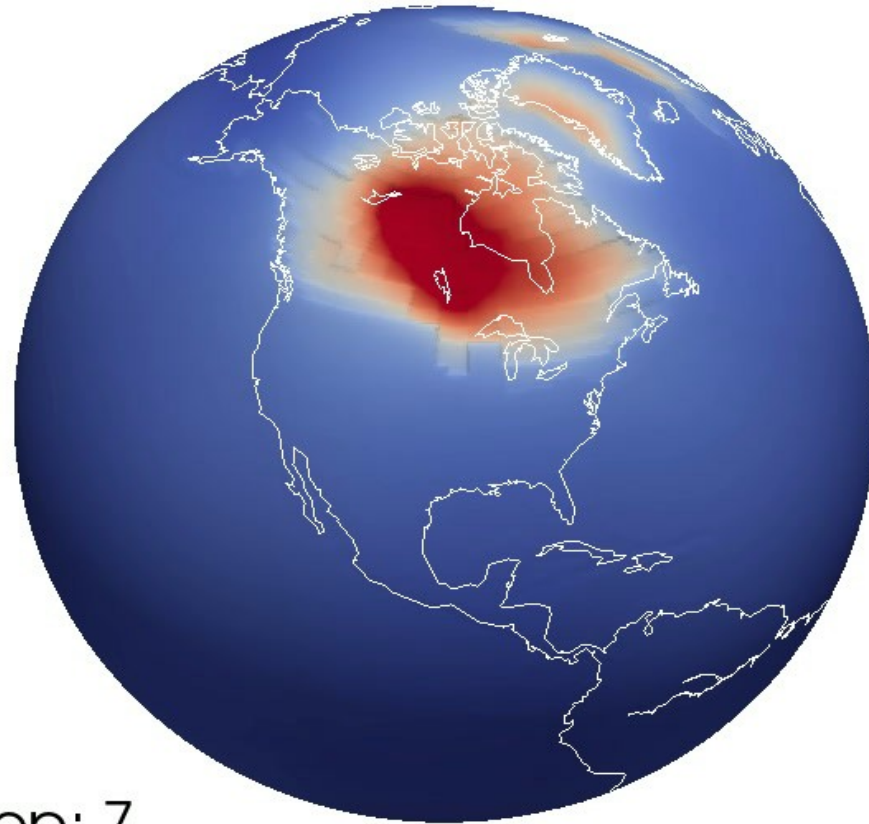
Infinite-element layer



Background gravity



Glacial rebound

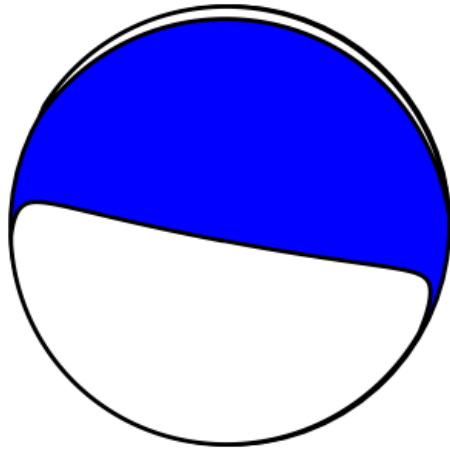


Time step: 7

Tidal loading

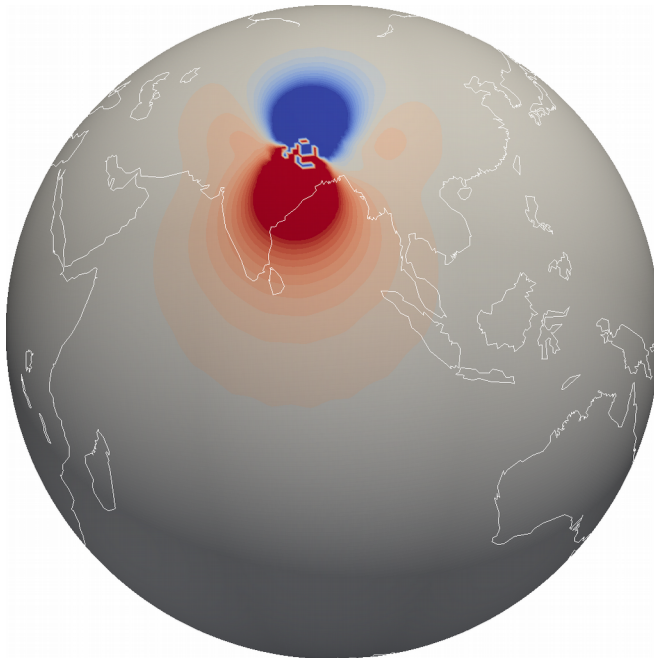


Postearthquake relaxation

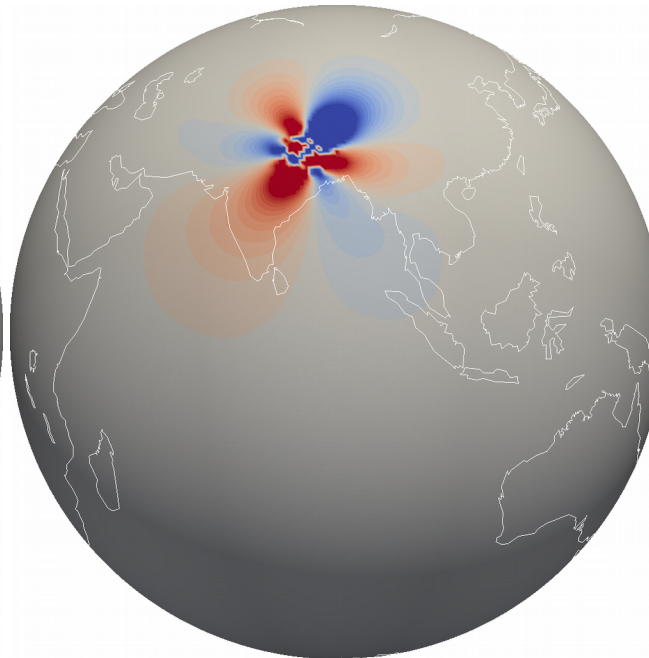


- Nepal Earthquake
- April 25, 2015
- 7.8 Mw

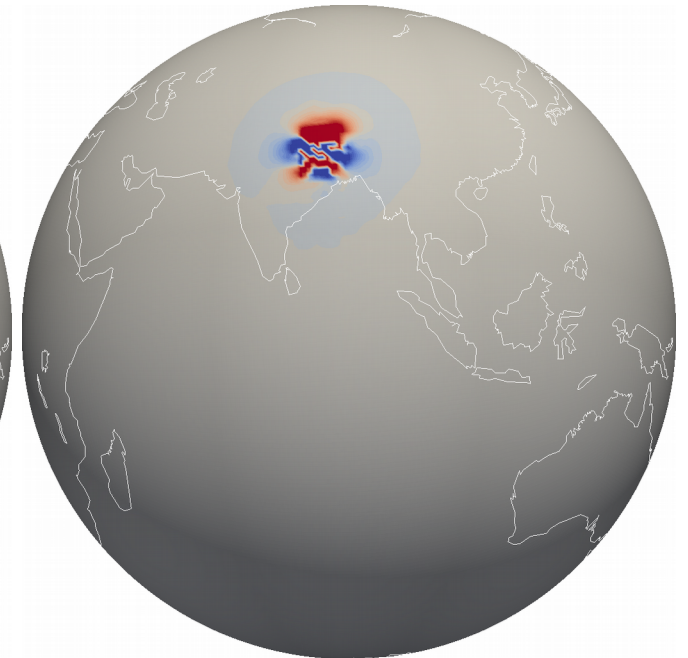
N



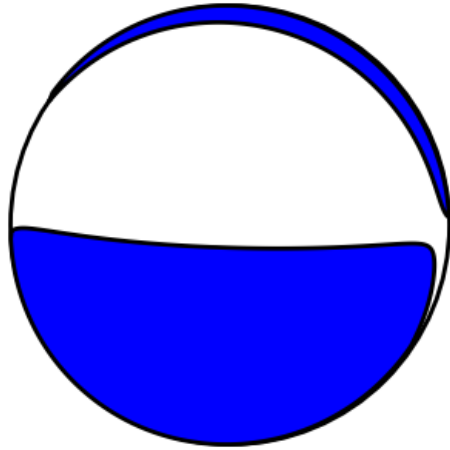
E



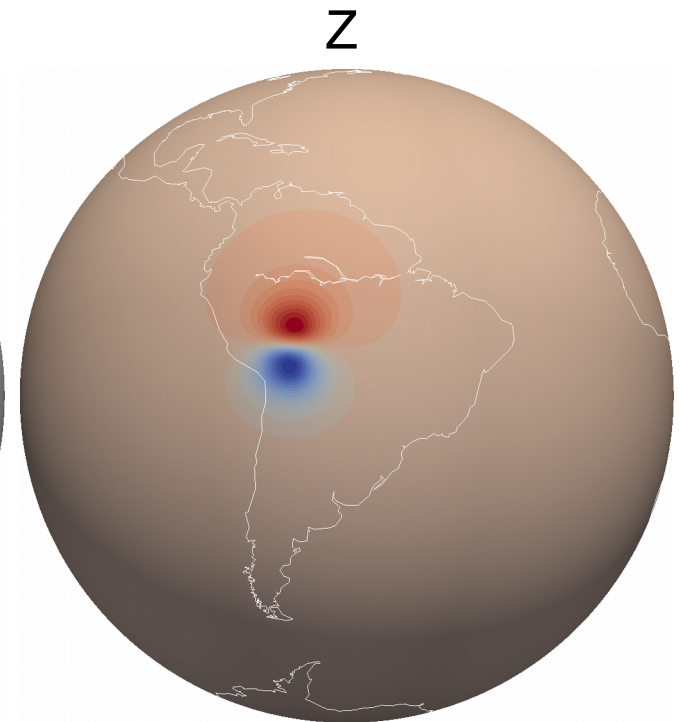
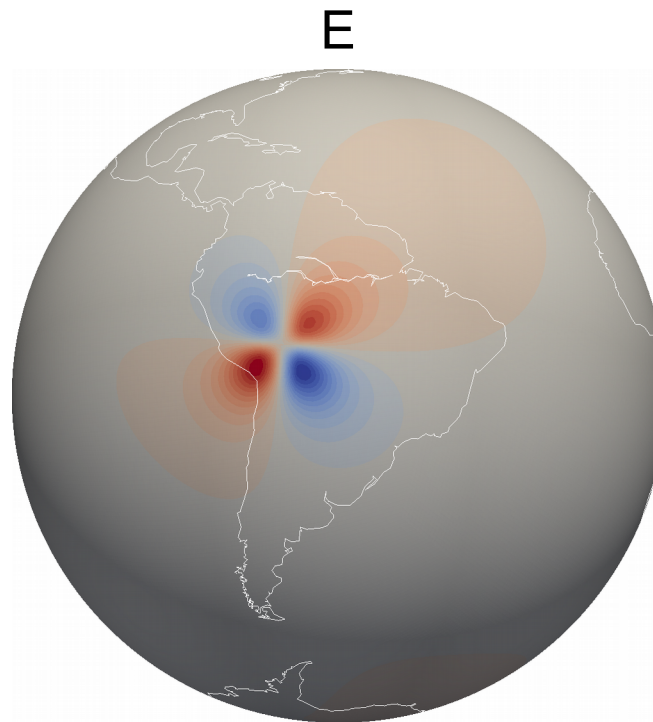
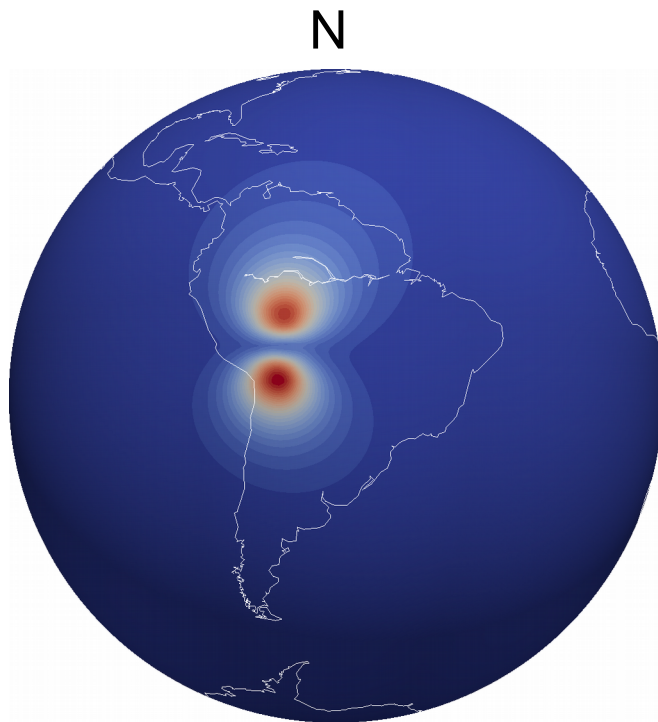
Z



Postearthquake relaxation

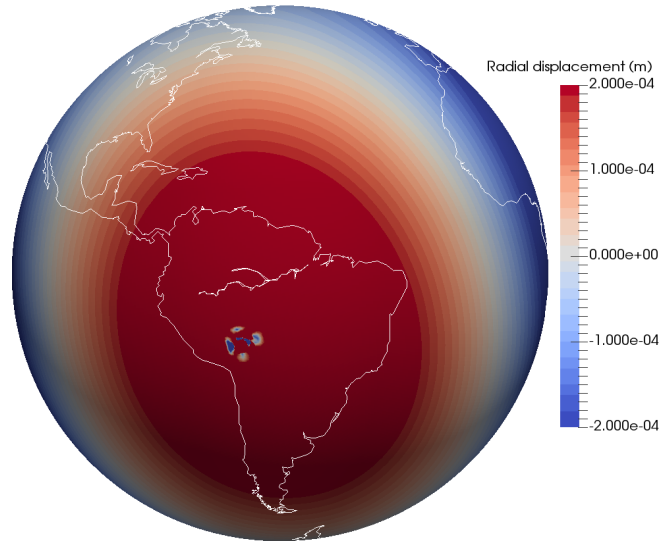


- Bolivia Earthquake
- June 09, 1994
- 8.2 Mw

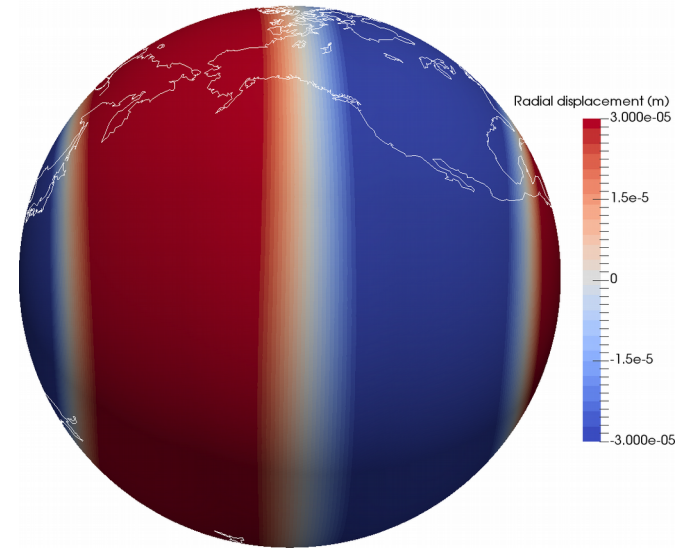


Frequency domain

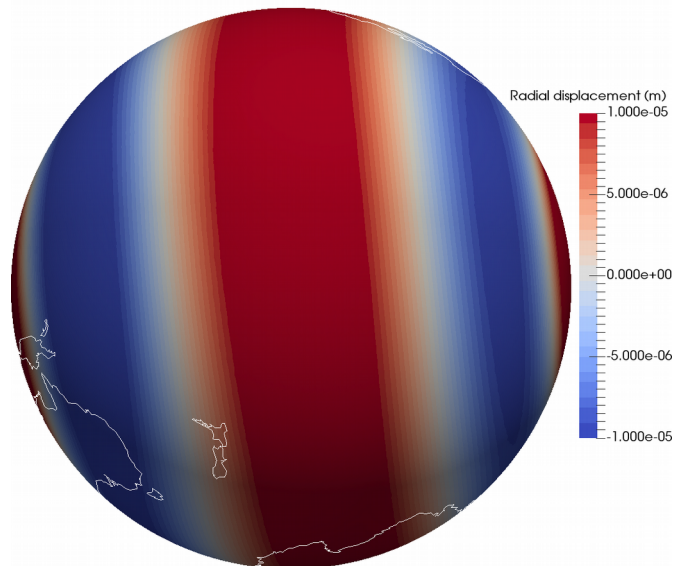
${}_0S_2$



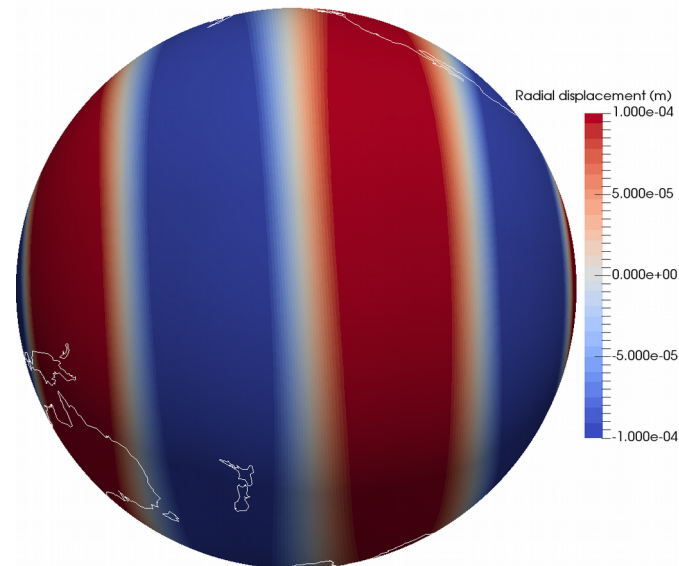
${}_0S_3$



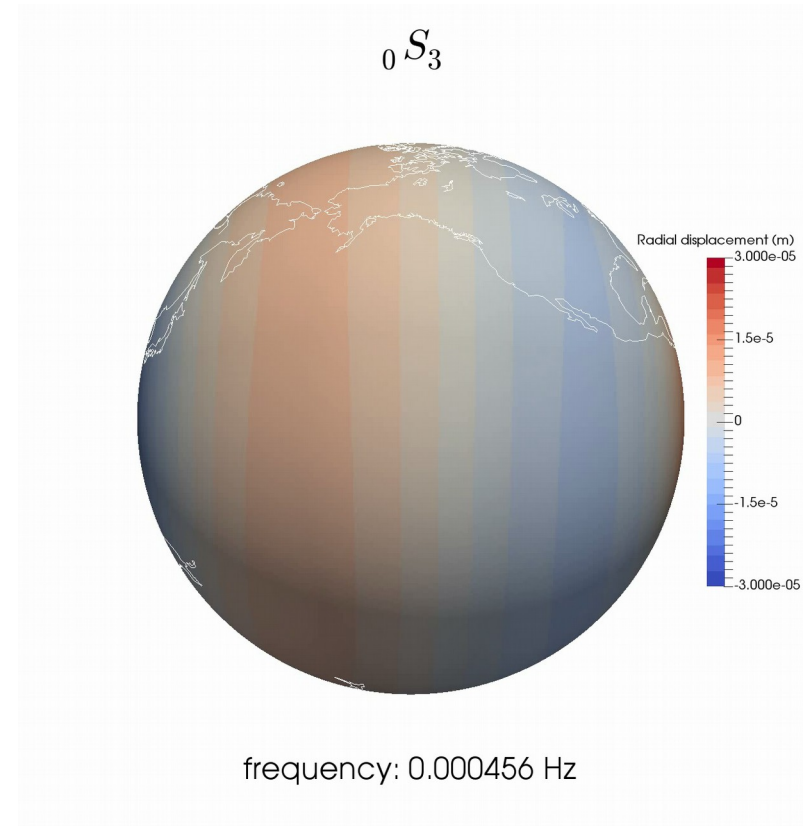
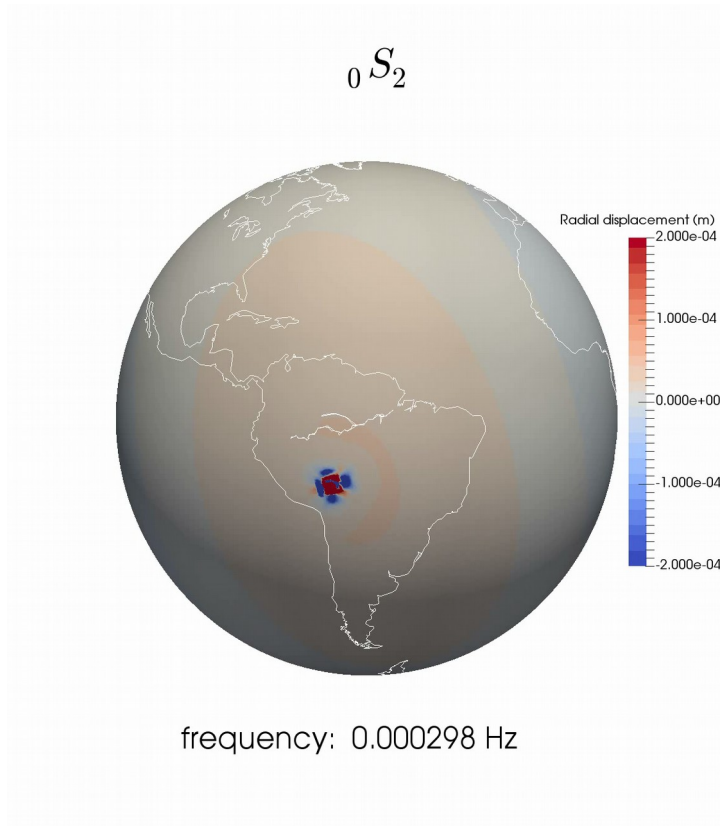
${}_0S_4$



${}_0S_5$



Frequency domain



Conclusion

- Solution of complete governing equations
- Applications
 - Wave propagation
 - Glacial isostatic adjustment
 - Normal modes
 - Tidal loading
 - Postearthquake relaxation
- Full 3D models